Jurnal Geologi Indonesia, Vol. 6 No. 1 Maret 2011: 51-68

Nannoplankton Assemblage Succession Throughout Cretaceous/ Tertiary Boundary in the "P" Well Section, Santos Basin, Brazil

Runtunan Kumpulan Nanoplankton pada batas Kapur/Tersier dalam Penampang Sumur "P", Cekungan Santos, Brasil

Panuju

PPPTMGB "LEMIGAS"
Jln. Ciledug Raya - Cipulir, Kebayoran Lama Jakarta Selatan - 12230

Abstract

The massive change in calcareous nannoplankton assemblages throughout Cretaceous/Tertiary (K/T) boundary (65.5 M.a.) has been illustrated by several authors. The diverse and abundant assemblage disappears suddenly above the Cretaceous/Tertiary boundary. This event is related to the most dramatic environmental changes in the Earth's history due to the catastrophic events, those are meteorite impact (*Chicxulub*) and supervolcano eruption (*Deccan*) occurring at the end of Cretaceous. The succeeding age was a time of rapid evolution of nannoplankton during Paleocene. A quantitative method analysis of nannoplankton throughout Maastrichtian to Paleocene of "P" well section, Santos Basin, Brazil, indicated that the nannoplankton assemblages abruptly decrease in diversity and abundance and mostly change in species composition. The various complex shapes of species at Maastrichtian also underwent changing to simple plain shapes and small at Paleocene. The sedimentary section ranges from the top of zone CC23 (Coccolith Cretaceous 23) to NP9 (Nannoplankton Paleogen 9). It is bounded by the Last Occurrence (LO) of *Tranolithus pachelosus* at the base and *Fasciculithus tympaniformis* at the top. The biostratigraphic discontinuity characterized by the absence of zone CC26 to NP4 is an indicator for the presence of an unconformity at K/T boundary within analyzed section. The Cretaceous nannoplankton assemblages are dominated by Genera *Watznaueria, Micula, Arkhangelskiella, Cribrosphaerella, Eiffellithus, Predicosphaera*, and *Retecapsa*, whilst the Paleocene assemblages are dominated by Genera *Toweius*, *Ericsonia*, and *Coccolithus.* Survivor Cretaceous species recovered into Tertiary sediments consist of *Braarudosphaera bigelowii, Biscutum melaniae, Neocrepidolithus neocrassus, Placozygus sigmoides, Cyclagelosaphaera reinhardtii, Markalius inversus,* and *Scapolithus fossilis*.

Keywords: biostratigraphy, nannoplankton, Cretaceous/Tertiary (K/T) boundary, Santos Basin, Brazil

Sari

Perubahan secara besar-besaran dalam kumpulan nanoplankton pada batas Kapur/Tersier (K/T) (65,5 jtl.) telah digambarkan oleh beberapa peneliti. Kumpulan yang beragam dan melimpah seketika menghilang pada batas Kapur/Tersier. Kejadian ini berhubungan dengan perubahan lingkungan paling dramatis dalam sejarah bumi yang disebabkan oleh bencana besar, yaitu meliputi tumbukan meteor (Chicxulub) dan letusan supervulkanik (Deccan) yang terjadi pada akhir Kapur. Periode berikutnya merupakan waktu evolusi nanoplankton yang berjalan secara cepat selama interval waktu Paleosen.Analisis dengan metode kuantitatif terhadap nanoplankton di sepanjang umur Maastrichtian sampai Paleosen dari penampang Sumur "P", Cekungan Santos, Brasil menunjukkan bahwa kumpulan nanoplankton seketika berkurang dalam keragaman dan kelimpahan serta mengalami perubahan besar dalam komposisi spesies. Bentuk spesies yang kompleks dan beragam pada Maastrichtian juga berubah menjadi sederhana dan kecil pada Paleosen. Penampang sedimen berumur antara puncak zona CC23 (Coccolith Cretaceous 23) sampai zona NP9 (Nannoplankton Paleogen 9).Umur tersebut diikat oleh kemunculan akhir (LO) Tranolithus pachelosus di bagian dasar dan Fasciculithus tympaniformis di bagian puncak. Ketidakselarasan pada batas Kapur/Tersier dijumpai, yaitu

Naskah diterima: 14 September 2009, revisi kesatu: 04 Desember 2009, revisi kedua: 02 Februari 2011, revisi terakhir: 11 Maret 2011

ditandai oleh ketidakhadiran zone CC26 sampai zona NP4. Kumpulan nanoplankton Kapur didominasi oleh genera Watznaueria, Micula, Arkhangelskiella, Cribrosphaerella, Eiffellithus, Predicosphaera, dan Retecapsa, sedangkan kumpulan Paleosen didominasi oleh Genera Toweius, Ericsonia, dan Coccolithus. Spesies Kapur yang bertahan dan dijumpai pada umur Tersier meliputi Braardosphaera bigelowii, Biscutum melaniae, Neocrepidolithus neocrassus, Placozygus sigmoides, Cyclagelosaphaera reinhardtii, Markalius inversus, dan Scapolithus fossilis.

Kata kunci: *biostratigrafi, nanoplankton, batas Kapur Tersier (K/T), Cekungan Santos, Brasil*

Introduction

Background

The massive change in calcareous nannoplankton assemblages at Cretaceous/Tertiary (K/T) boundary was first noted and illustrated by Bramlette & Martini (1964) and then described in greater detail by Perch-Nielsen (1969, 1979a-b, 1981), Percival & Fischer, 1977, Romein (1977), and Bown (1999). The diverse and abundant Maastrichtian assemblage disappears suddenly at the K/T boundary. It is then replaced by new species and genera evolving from 15 to 18 genera that survived the K/T boundary event (Perch-Nielsen, 1982). However, some survivors are considered as reworked forms by some authors since their occurrence in Tertiary are always very rare, even in the abundant assemblage. The confirmed data reveal that survivors consist of some 12 species of 11 genera from 10 families. The succeeding age was a time of rapid evolution of nannoplankton. There was evident that about 25 new genera occurred during ±12.5 M.a. of Paleocene time interval (Perch-Nielsen, 1985) that evolved from 10 survivors (Bown, 1999).

In this paper, the kind of nannoplankton succession throughout Cretaceous/Tertiary boundary can be seen on the “P” well section in Santos Basin (Figure 1). This basin is included in the basins having the ideal sedimentary section to identify the succession marine organism during K/T boundary. Identification of the succession done is restricted on the age of Late Maastrichtian to Early Paleocene.

Material and Methods

This study is the result of nannoplankton analysis on the 27 samples from “P” well section comprising ditch cuttings and cores. They are processed mainly using *smearing method* and embedded in entellan. The analyzed interval was determined systematically and the observation was undertaken at a magnification of 1000x using a light microscope (LM) in a quantitative method. Observation techniques comprise *bright field (BF), cross polarized light (XPL), Gypsum plate in XPL,* and *phase contrast*. Taxonomy and terminology in the description of index species refers to Perch-Nielsen (1985). The standard zonation of Sissingh (1977) and Martini (1971) is used as a mainframe for biostratigraphic subdivision, and then the result is used to identify the succession of nannoplankton assemblage throughout Cretaceous/Tertiary boundary. The flow chart of research method can be seen in Figure 2.

Cretaceous/Tertiary Boundary

The K/T boundary that marks the separation between Cretaceous and Tertiary is visible in the geological record by a discontinuity (dramatic change) in the fossil development. This boundary corresponds to one of the greatest mass extinctions in Earth history. At least 75 percent of the species on our planet, both in the seas and on the continents, were extinguished. In the oceans, more than 90 percent of the plankton was extinguished, which inevitably led to the collapse of the oceanic food chain (Figure 3). All ammonites, genuine belemnites, and rudistids are extinct, and most species of foraminifera and nannoplankton, diatoms,dinoflagellates, molluscs, echinoids, fish, and marine reptiles disappeared. Even though some groups, such as squids, octopus, nautilus, and a few species of foraminifera and nannoplankton, diatoms, dinoflagellates, brachiopods, molluscs, and echinoids survived, the genetic pool were relatively very small at the dawn of the Tertiary Period. The recovery of the marine biota after K/T boundary event was fairly rapid after the mid-Paleocene due to overall transgressing seas and ameliorating cli-

Figure 1. Studied area of Santos Basin (Modified from HIS., 2007).

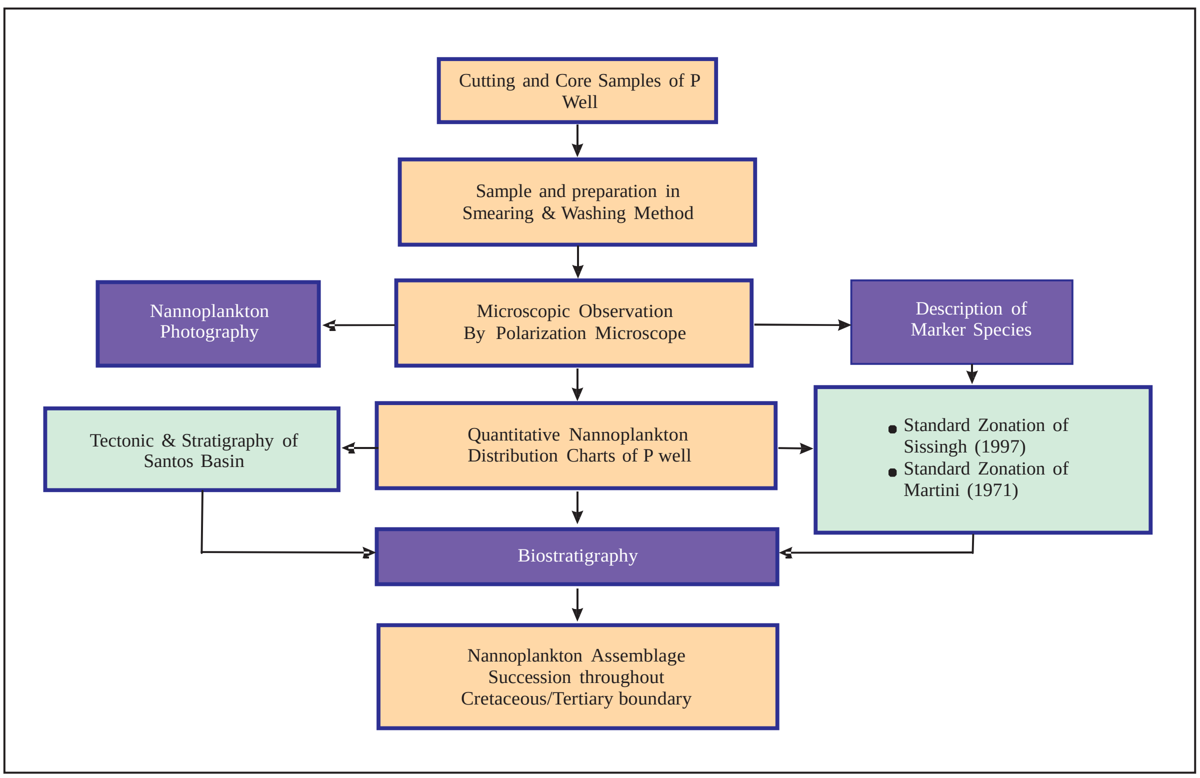


Figure 2. Flow chart of the research method.

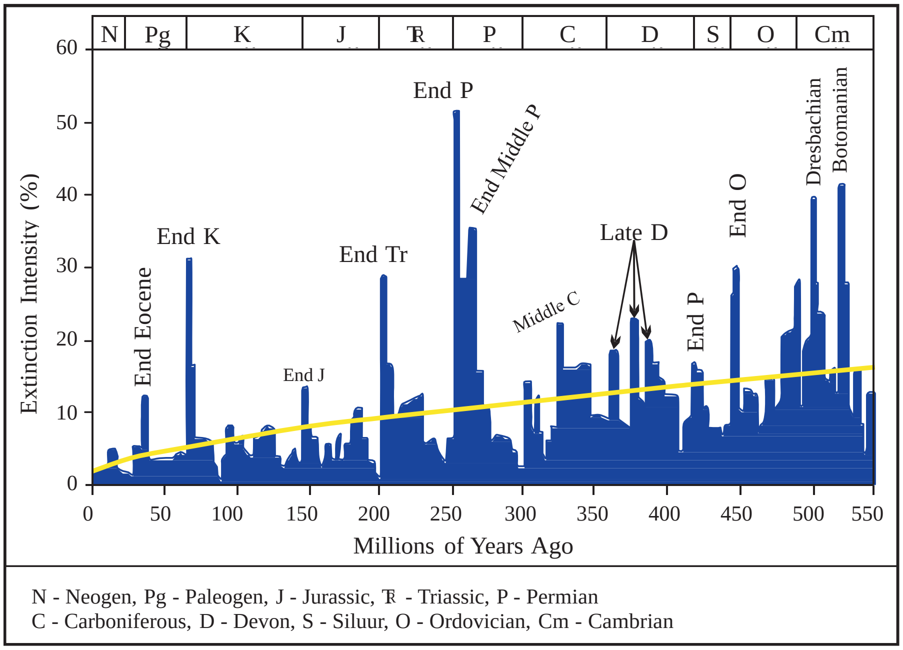


Figure 3. Extinction intensity of marine genus biodiversity (After Raup & Sepkowsky, 1977 *in* Wikipedia, 2007).

mates. On the continent, the large dinosaurs which had been decline for over 20 million years, died out forever. However, most mammals, birds, turtles, crocodiles, lizards, snakes, amphibians, and some land flora were primarily unaffected by the End-Cretaceous mass extinction (Raup & Sapkowsky 1977, *in* Wikipedia, 2007).

There are debates about causes of mass extinction in K/T boundary, and the explanation is present within sedimentary rocks. Rocks deposited during the Cretaceous and Tertiary Periods are separated by thin clay layers that are visible at several sites around the world (Figure 4). A team of scientists led by Alvarez *et al.* (1980) discovered the clay layers contains strikingly high concentration of iridium, an element that is much more common in meteorites (asteroid or comet) than in Earth crustal rocks. Consequently, they suggested that the meteorite impacts

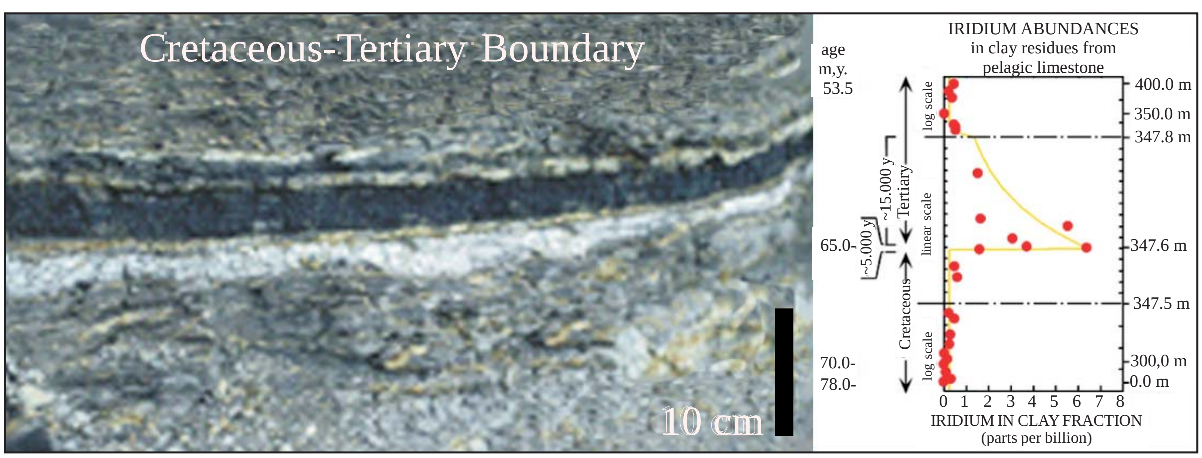


Figure 4. The thin clay layer contains a strikingly high concentration of iridium between K/T boundary (Modification Sharpton *et al.*, 1992).

have generated the iridium anomaly (Sharpton *et al.*, 1992). This argument is supported by the discovery of shocked quartz (Figure 5), microspherules, and mega wave deposits. The high iridium concentrations in the clay layers at several places around the world suggested the impact was a large one. Eventually, most paleontologists began to accept the idea that the mass extinctions at the end of the Cretaceous were largely or at least partly due to a massive meteorite impact (Wikipedia, 2007). The most famous evident for meteorite impact is the 180 km diameter of the buried *Chicxulub Crater*, on the coast of Yucatan, Mexico (Figure 6). Even, some scientists conclude, that the mass extinction event during K/T boundary is not caused by a single impact since multiple impacts appear to be very common

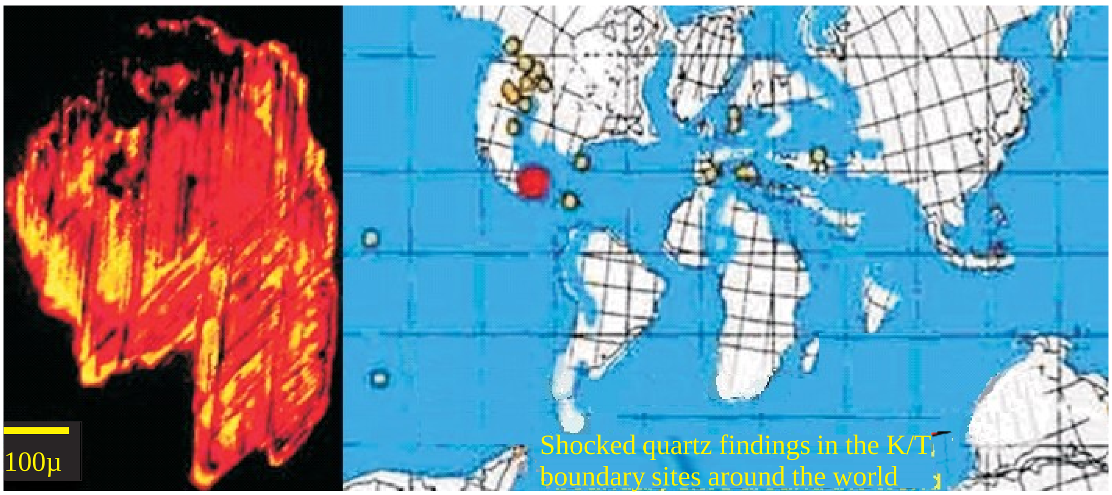


Figure 5. Shocked quartz and its distribution (Modification from Sharpton *et al.*, 1992).

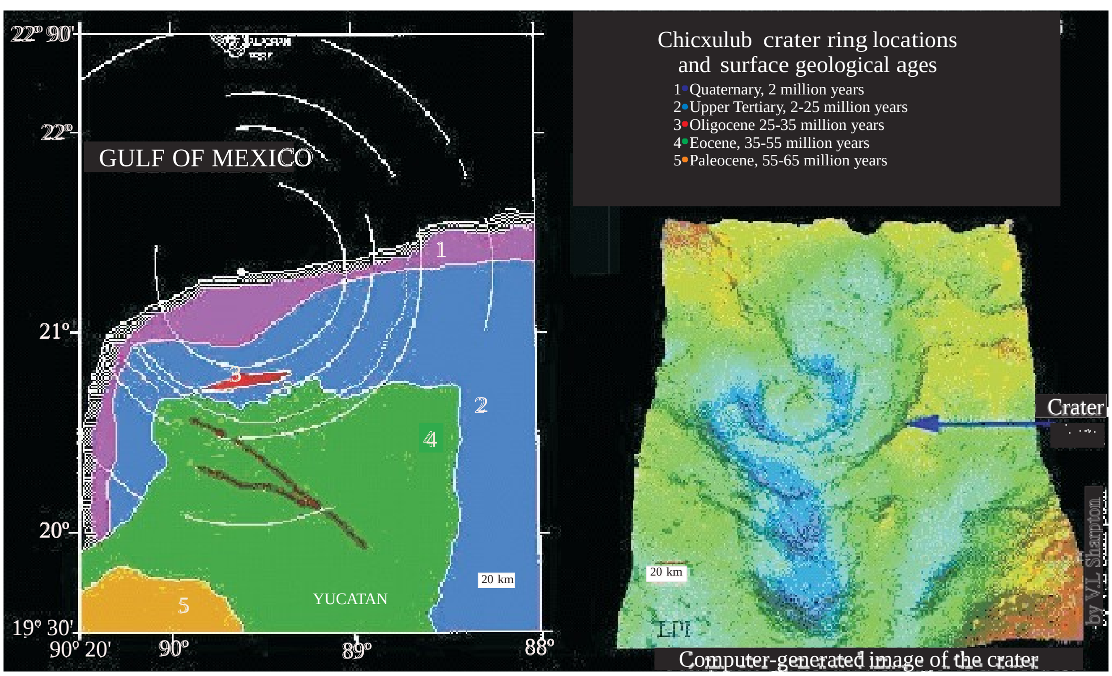


Figure 6. Chicxulub Crater, Yucatan, Mexico (Modification from Sharpton *et al.*, 1992).

throughout the solar system. The Shiva crater is another huge impact crater located under the Arabian Sea off the coast of India near Bombay. This crater also dates from the K/T boundary, 65 M.a., when the Chicxulub crater at the tip of the Yucatán Peninsula also formed. Although it has shifted because of sea floor spreading, when pieced together, it would be about 370 miles (600 km) by 280 miles (450 km) across and 7.5 miles (12 km) deep (and may be just part of a larger crater). It is estimated to have been made by a bolide (an asteroid or meteoroid) 25 miles (40 km) in diameter. This crater was named by the paleontologist Sankar Chatterjee for Shiva, the Hindu god of destruction and renewal. The other craters are Boltysh crater (24 km diameter, 65.17 ± 0.64 M.a.) in Ukraine, Silverpit crater (20 km diameter, 60-65 M.a.) in the North Sea, Eagle Butte crater (10 km diameter, < 65 M.a.) in Alberta, Canada, and Vista Alegre crater (9.5 km diameter, < 65 M.a.) in Paraná State, Brazil (Wikipedia, 2007.). The nemesis hypothesis of Raup and Sepkowsky, 1977 (*in* Wikipedia, 2007) theorizes that there is a periodicity of 26 million years to mass extinctions which is caused by collisions with a comet from the Oor cloud as they are perturbed in their orbits by a dark star (a companion star to the sun) (Geolor's Earth Issues, 2007).

Some scientists considered that there was a link between large impacts and volcanic eruptions. This is evidenced by *Deccan Trap* in India during K/T boundary (the second greatest volcanic eruption after *Siberian Trap* during Permian/Triassic (P/ TR) boundary). By some scientists, The *Deccan Traps* is also assumed as an agent which had been contributed to the extinction in the end of Cretaceous (Wikipedia, 2007).

Santos Basin

The Santos Basin covers an area of 352,260 km2, bordered by Florianopolis High (Pelotas) and Cabo Frio High (Campos). As part of the rifted Atlantic margin of South America, the geological history of the Santos Basin can be divided into pre-Rift (pre-Cretaceous), syn-Rift (Neocomian to Barremian), and post-Rift (Aptian to Recent) stages, as shown on the seismic-based interpretation in Figure 7 (Joshua, 2007).

The lithology and age of pre-Rift rocks in the deepwater Santos Basin are opened to speculation. Reassembly of the African and South American cratons suggests that the Santos' pre-Rift units are a crystalline complex that may contain late stage pre-Rift basaltic flows and intrusions. Field modeling based on interpretation of the Veritas 3D seismic data, plus the fact that this basin is adjacent to the Campos Basin, suggest that the syn-Rift units probably comprise lacustrine, continental, and neritic facies with possible basaltic intrusions and lava flows (Bagni, 2007).

In much of the deepwater area, the basal post-Rift unit is Aptian Salt of more than 2,000 m thick, which forms an excellent seal for the syn-Rift sequence related to the petroleum system. Overlying the salt are Albian deepwater carbonates and marls, which are overlain in turn by Late Cretaceous through Recent turbiditic clastics. The thickness of these clastics varies within basin that floored by salt and/or salt welds, but in general, the clastics thin seaward (Joshua, 2007).

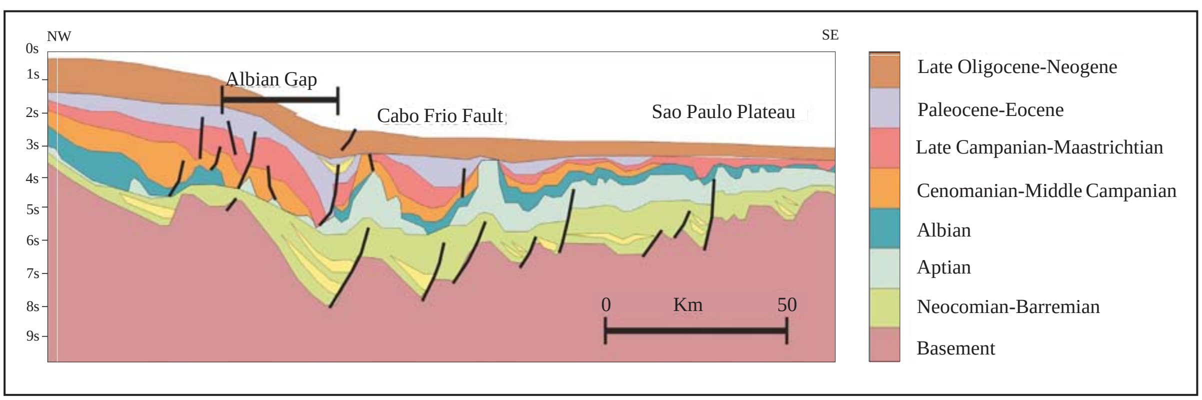


Figure 7. The 2D seismic-based interpretation across the deepwater Santos Basin (Joshua, 2007).

The tectonic-stratigraphic evolution was commenced by Mesozoic pelotas to PE-PB rift followed by the occurrence of Neocomian basalt of the Camboriú Formation. Basin Rifting continued during Barremian/Early Aptian sequence with the result of Guaratiba Formation. The transitional stage is evidenced by the formation of Aptian Evaporites (Ariri Formation). Entering the Albian, the drifting occurred with the result Guaraja Formation (Bagni, 2007).

The further sequences are the results of the Cenomanian-Turonian transgressive phase (Itajai-Açu Formation), the Coniacian to Maastrichtian regressive phase (Juréia Formation) due to Serra do Mar Uplifting, and finally by Tertiary Transgressive phase (Iguape and Marambaia Formations) (Bagni, 2007). Based on a seismic stratigraphic analysis, three major sequences from their internal reflector patterns and external geometry in the southern end of the basin can be identified. The Early Rift Sequence (lower rift on Tupi seismic line) is compounded by volcanic rocks and characterized by parallel to subparallel reflectors, continuous and high dip angles. "The Rift Sequence (upper rift on Tupi seismic line) is characterized by half grabens, possibly filled by coarse sediments. The main internal reflections are divergent and prograding. The final Sag Phase was deposited on an unconformity identified by 3-D data. Reflectors truncate down and onlap above (Joshua, 2007). Regional stratigraphic chart of Santos Basin can be seen in Figure 8.

Analysis Results

Biostratigraphic Subdivision

The standard zonation of Sissingh (1977) for Cretaceous and Martini (1971) for Tertiary are used in the age interpretation and zonal subdivision throughout “P” well section (Table 1). Based on the occurrence of nannoplankton markers, biostratigraphy of the studied section can be subdivided from the base to the top as follows (Figure 9a-f).

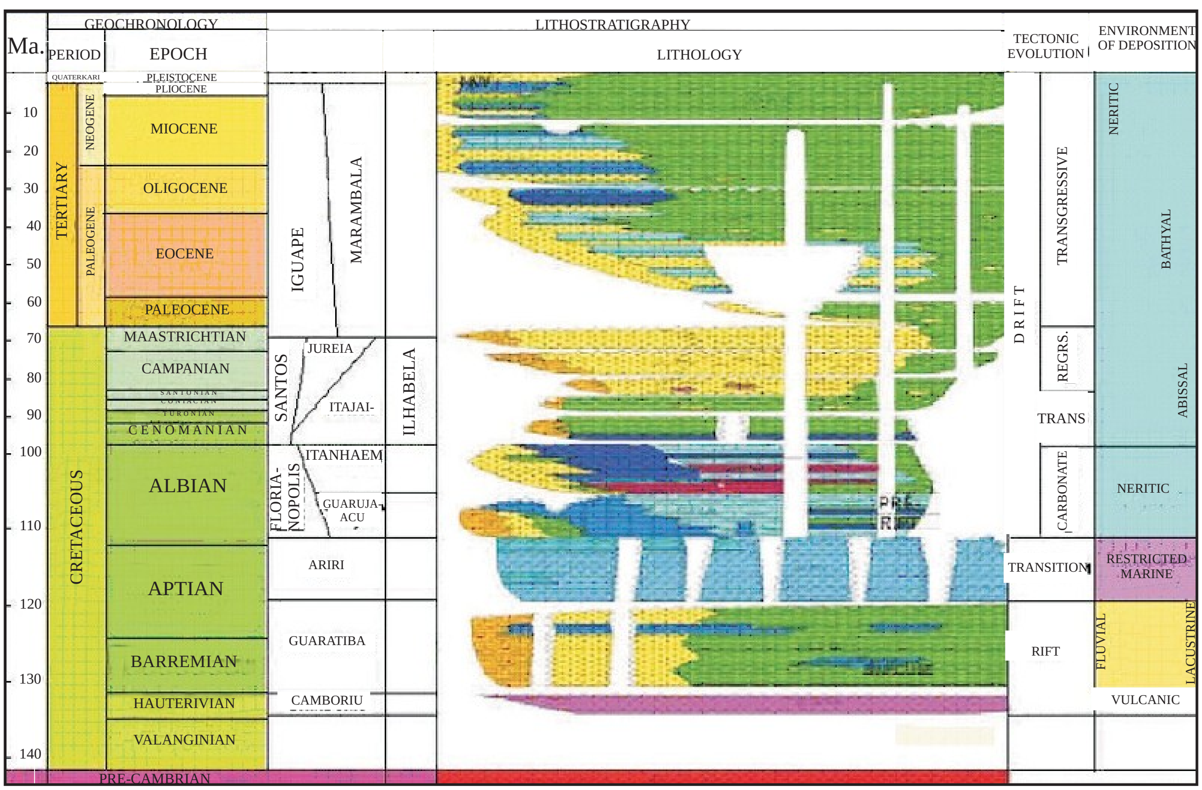


Figure 8. Regional stratigraphy of the Santos Basin (Bagni, 2007).

Table 1. Biostratigraphy of P Well Section

Sample	Zone	Description
17-27	Zone NP9 (Late Paleocene)	This zone is bounded respectively by the FO of *Discoaster multiradiatus* at the base and the occurrence of *Fasciculithus tympaniformis* at the top. This zone can be subdivided into subzone "a" and "b" by the FO of *Campilosphaera eodela* in sample 21.
8-16	Zone NP8 (Late Paleocene)	This zone is bordered by the FO of *Heliolithus riedelii* at the base and the FO of *Discoaster multiradiatus* at the top.
7	Zone NP7 (Late Paleocene)	The bottom of this zone is revealed by the FO of *Discoaster mohleri*, whilst its top is marked by the FO of *Heliolithus riedelii*.
5-6	Zone NP6 (Middle Paleocene)	The base and the top of this zone are indicated respectively by the FO of *Chiasmolithus bidens* and the FO of *Discoaster mohleri.*
4	Zone NP5 (Middle Paleocene)	The base of this zone is underlain by an unconformity ranges from CC26 to NP4 (uppermost part of Maastrichtian to lower part of Middle Paleocene). This zone is bounded by the occurrence of *Fasciculithus janii* at the bottom and the first occurrence (FO) of *Chiasmolithus bidens* at the top.
3	Zone CC25 (Maastrichtian)	The bottom of this zone is assigned by the LO of *Reinhardtites levis*, whilst its top is indicated by the occurrence of *Calculites obscurus*. This zone is deposited at the top of Cretaceus sedimentary sequence of "P" well section. Zone CC26 is considered to be absent since the LO of *Calculites obscurus* coincided with the highest occurrence of species ranges to CC26. This is supported by the absence of species that restricted within zone CC26 *(Ceratolithoides kamptneri, Nephrolithus frequens, Cribosphaera daniae* and *Micula prinsii)* and the occurrence of Middle Paleocene assemblage within the overlying sample (4).
2	Zone CC24 (Maastrichtian)	This zone is indicated by the LO of *Tranolithus pachelosus* at the base and the LO of *Reinhardtites levis* at the top.
1	Zone CC23 (Maastrichtian)	This sample is the position of the last occurrence (LO) of *Tranolithus pachelosus* that indicates the top of zone CC23.

Nannoplankton Assemblage Succession throughout Cretaceous/Tertiary Boundary in the "P" Well Section

A quantitative method nannoplankton analysis throughout the top of zone CC23 (Maastrichtian) to zone NP9 (Paleocene) of "P" well section, Santos Basin, Brazil has been undertaken to define the succession of nannoplankton assemblage throughout K/T boundary. Unfortunately, the real succession within K/T boundary cannot be seen due to the presence of an unconformity ranges from CC26 to NP4 (Figures 9a-f & 10). However, it was highly evident that the nannoplankton assemblages abruptly decreased in diversity and abundance and mostly changed in species composition (Figure 11). Mass extinction had been occurred in the Cretaceous nannoplankton assemblage during K/T boundary event and only 7 survivors can be recovered (Figure 11a-d). The domination of large various complex shapes of species at the upper part of Maastrichtian

AGE (Ma)	PERIODS	EPOCH	STAGE	ZONE: CC (Sissingh, 1977); NP (Martini, 1971)		BIOMARKERS	DEPTH (Represented by Number)	Fasciculithus janii	Coccolithus pelagicus	Chiasmolithus consuetus	Cruciplacolithus tenuis	Ellipsolithus macellus	Ericsonia robusta	Ericsonia subpertusa	Ericsonia spp.	Sphenolithus primus	Prinsius bisulcus	Fasciculithus tympaniformis	Toweius selandianus	Toweius spp	Prinsius sp.	Fasciculithus sp	Toweius pertusus	Chiasmolithus bidens	Discoaster multiradiatus	Fasciculithus alanii	Fasciculithus lilinae	Ellipsolithus distichus	Sphenolithus sp	Semihololithus kerabyii	Fasciculithus clinatus	Heliolithus cantabriae	Fasciculithus involutus	Toweius rotundus	Zygodiscus herlynii	Sphenolithus primus	Sphenolithus anarrhopus	Toweius eminens	Campylosphaera eodela	Lophodolithus nascens	Sphenolithus small	[illegible]
55.8	TERTIARY	PALEOCENE LATE	THANETIAN	NP9	b	Fasciculithus tympaniformis	27		230		2	2	7	5	5	17	2	7	17	35		4	58	1	2	6	1	2	2	2	2	1	1	175	1	7	4	1	2	2	5	1
							26		43			1	2	1	2	23		5	5	3		2	9	1		1			4	1	3			35	2	9	2		1	1	7	1
							25		40	3	1	1	5	2	8	11	2	4	14	2		2	51			3		1	3		1		1	125		6	5	2	1	1	2	
							24		52	2		2	1	2		11		1	2	1		1	7	1		1		1	3					41		3	2		1			
							23		315	6	2	7	85	72	118	54	5	3	46	34		1	125	2	1		1	2	11		3			435	2	5	35	3	8	2		4
							22		295	7	3	6	63	43	102	21	10	5	86	43		2	168	2	1				5		1			780	1	2	17		4	1		5
						Campilosphaera eodela	21		375	7	5	21	69	64	175	23	3	13	17	31		2	142		2		3	3	7		2			1080	1	2	13	5	7	7		3
					a		20		7				1	2		3	2	2	8	3		1	2			1			2					4			2					
							19		6		1		1	2		3	2	2	6	2			7	1					1					5		1	1					
							18		23	1	1	4	2	5	11	27	11	7	125	18		4	17	1		1	1	1	14		1		2	11		3	11	1				1
						Discoaster multiradiatus	17		265	3	1	3	55	43	75	65	15	13	175	23		7	35	2	2	1	1	1	27		2		3	425	1	5	17					2
				Np8			16		5							4	1		5	3			12						3					2	1							
							15		1			1				2			2	1									2													
							14		2				3	1	3	2													2		1											
							13		1			1	1		1	2		1	1										1													
							12		1		1				3					1									3				1									
							11		2		1				2	1	2	1	6	1								1	2			1										
							10		2		3				6	2	1		3	2									3													
							9		1	1	1				2		1		1																							
						Heliolithus riedelii	8		5	1	1	1	2	2	3	1	2	2	7	1			2	1				1	1		2											
				NP7		Discoaster mohleri	7	1	8	1	3	2	8	10	7	7	5	2	21	2		1						3	8	1												
58.7				NP6			6		1					1		1		1	1																							
		MIDDLE	SELAND.			Chiasmolithus bidens	5	1	2		1			1	1			2				1	1	1																		
63.0				NP5		Fasciculithus janii	4	2	2	1	2	1	1	2	1	1	1	1	2	2	1																					
66,5	CRETACEOUS	LATE	MAASTRICHTIAN	CC25		Calculites obscurus	3																																			
				CC24		Reinhardtites levis	2																																			
70,0				CC23		Tranolithus pachelosus	1																																			

LEGENDS: Last Occurrence, First Occurrence, Occurrence, R: Reworked forms

Next

Figure 9a. Quantitative Paleocene nannoplankton distribution chart and biostratigraphy of “P” Well Section.

AGE (Ma)	PERIODS	EPOCH		STAGE
55.8 – 58.7	TERTIARY	PALEOCENE	LATE	THANETIAN
58.7 – 63.0	TERTIARY	PALEOCENE	MIDDLE	SELAND.
66,5 – 70,0	CRETACEOUS	LATE		MAASTRICHTIAN

DEPTH (Represented by Number)	ZONE: CC (Sissingh, 1977); NP (Martini, 1971)	BIOMARKERS	Discoaster falcatus	Lophodolithus sp-x	Semihololithus biskayae	Discoaster binodosus	Fasciculithus thomasii	Cruciplacolithus frequens	Zygodiscus bramlettei	Indeterminate nannolith	Chiasmolithus sp.	Fascichulithus schaubi	Pontosphaera spp.	Discoaster mohleri	Discoaster nobilis	Discoaster falcatus	Discoaster limbatus	Neochiastozygus chiastus	Neochiastozygus distentus	Neochiastozygus sp.	Pontosphaera desueta	Neocrepidolithus biskayae	Neocrepidolithus bukiyi	Neochiastozygus junctus	Pontosphara multipora	Rhabdosphaera spp	Discoaster araneus	Zygodiscus plectopons	Scapolithus rhombiformis	Micrantolithus vesper	Micula decussata	Neochiastozygus concinnus	Neochiastozygus perfectus	Heliolithus nedelli	Cruciplacolithus latipons	Elipaolithus bollii	Fasciculithus bollii
27	NP9 b	Fasciculithus tympaniformis	1	1	1																																
26	NP9 b					1	1	1	1																												
25	NP9 b						2	2		1	2	1																									
24	NP9 b						1				1		1																								
23	NP9 b					2	1				3		23	2	1	1	1	5	7	2	1	6	7	4	1	1	1										
22	NP9 b					2	3				3		10	1		1		4	10	1	1	5	4	6				1	1								
21	NP9 b	Campilosphaera eodela				3	7				2		21	2	1	2	1	56	35	5	1	7	5	7	1	2		3		2	1						
20	NP9 a																																				
19	NP9 a																																				
18	NP9 a					1	1							1		1																					
17	Np8	Discoaster multiradiatus																																			
16	Np8																																				
15	Np8																																				
14	Np8																																				
13	Np8																																				
12	Np8																																				
11	Np8																																				
10	Np8																																				
9	Np8																																				
8	Np8	Heliolithus riedelii																																			
7	NP7	Discoaster mohleri																																			
6	NP6																																				
5	NP6	Chiasmolithus bidens																																			
.	NP5	Fasciculithus janii																																			
3	CC25	Calculites obscurus																																			
2	CC24	Reinhardtites levis																																			
1	CC23	Tranolithus pachelosus																																			

Prev. Next

LEGENDS: Last Occurrence First Occurrence Occurrence R: Reworked forms

Figure 9b. Quantitative Paleocene nannoplankton distribution chart and biostratigraphy of “P” Well Section.

AGE (Ma)	PERIODS	EPOCH	STAGE	ZONE: CC (Sissingh, 1977); NP (Martini, 1971)	BIOMARKERS	DEPTH (Represented by Number)	Braarudosphaera bigelowii	Markalius inversus	Scapolitus fossilis	Placozygus sigmoides	Biscutum melaniae	Neocrepidolithus neocarassus	Cyclagelosphaera reindhardtii
55.8	TERTIARY	PALEOCENE LATE	THANETIAN	NP9 b	Fasciculithus tympaniformis	27	1	1	1				
						26		1					
						25	1						
						24							
						23	3	2	3	3			
						22	2	1	2	5			
				NP9 a	Campilosphaera eodela	21	3	2	4	7			
						20	2						
						19							
						18	1	1		1			
				Np8	Discoaster multiradiatus	17	1	1	2	2			
						16	1		1				
						15	1						
						14	1						
						13			1				
						12	1	1					
						11		1					
						10	1						
						9							
				NP7	Heliolithus riedelii	8	1		1	1	1		
				NP6	Discoaster mohleri	7	1		1	1	2		
58.7		MIDDLE	SELAND.			6							
				NP5	Chiasmolithus bidens	5		1					
63.0					Fasciculithus janii	4	1					1	1
66,5	CRETACEOUS	LATE	MAASTRICHTIAN	CC25	Calculites obscurus	3	2	2	1	2	13	1	
				CC24	Reinhardtites levis	2	1		1	1			
70,0				CC23	Tranolithus pachelosus	1	2	1		1	2	1	1

LEGENDS: Last Occurrence; First Occurrence; Occurrence; R: Reworked forms

Prev. Next

Figure 9c. Quantitative Survivor nannoplankton distribution chart and biostratigraphy of "P" Well Section.

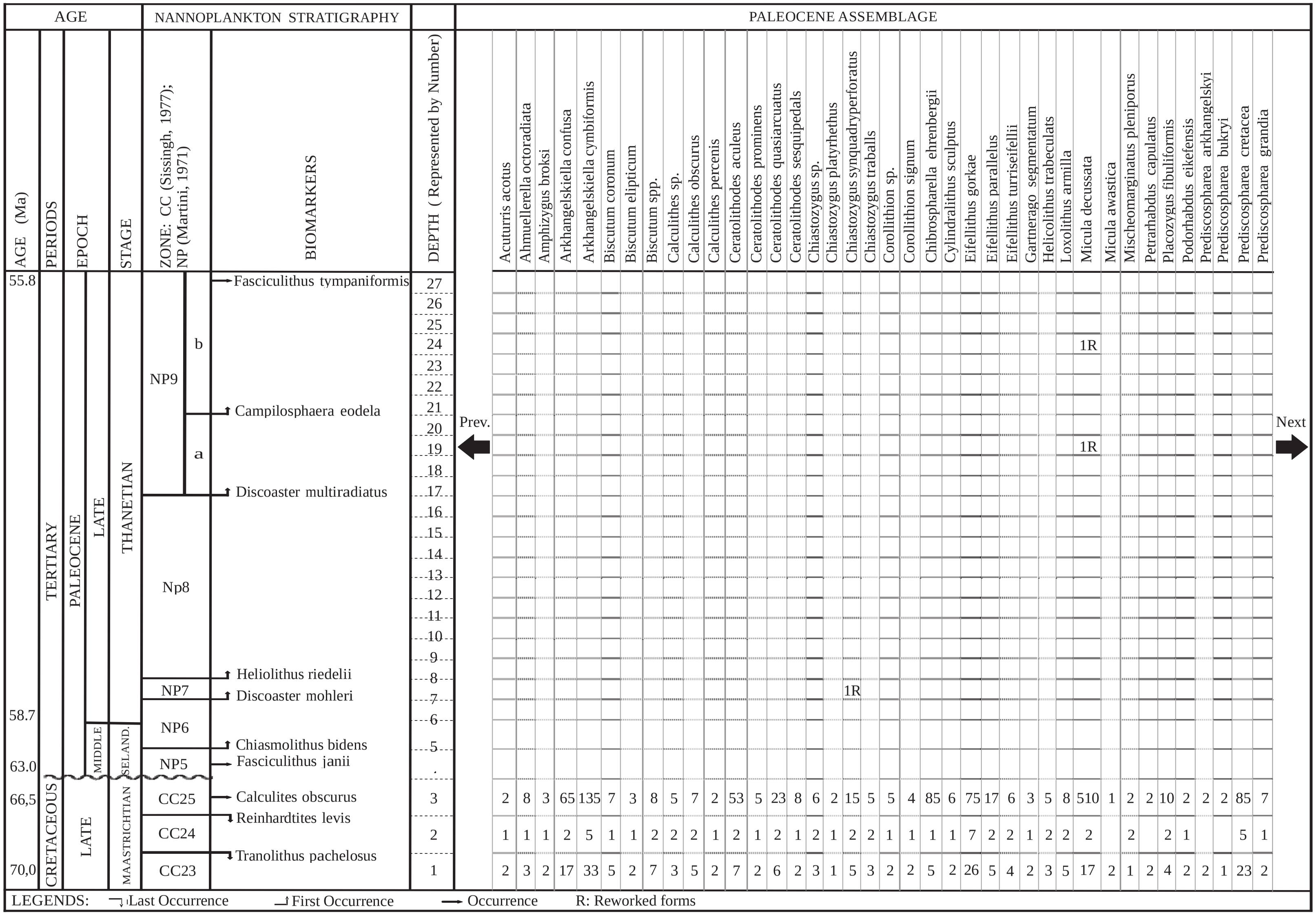


Figure 9d. Quantitative Cretaceous nannoplankton distribution chart and biostratigraphy of “P” Well Section.

AGE (Ma)	PERIODS	EPOCH	STAGE	ZONE: CC (Sissingh, 1977); NP (Martini, 1971)	BIOMARKERS	DEPTH (Represented by Number)
55.8	TERTIARY	PALEOCENE, LATE	THANETIAN	NP9 b	Fasciculithus tympaniformis	27
						26
						25
						24
						23
						22
				NP9 a	Campilosphaera eodela	21
						20
						19
						18
				Np8	Discoaster multiradiatus	17
						16
						15
						14
						13
						12
						11
						10
						9
				NP7	Heliolithus riedelii	8
				NP6	Discoaster mohleri	7
58.7		PALEOCENE, MIDDLE	SELAND.			6
				NP5	Chiasmolithus bidens; Fasciculithus janii	5
63.0						.
66,5	CRETACEOUS	LATE	MAASTRICHTIAN	CC25	Calculites obscurus; Reinhardtites levis	3
				CC24	Tranolithus pachelosus	2
70,0				CC23		1

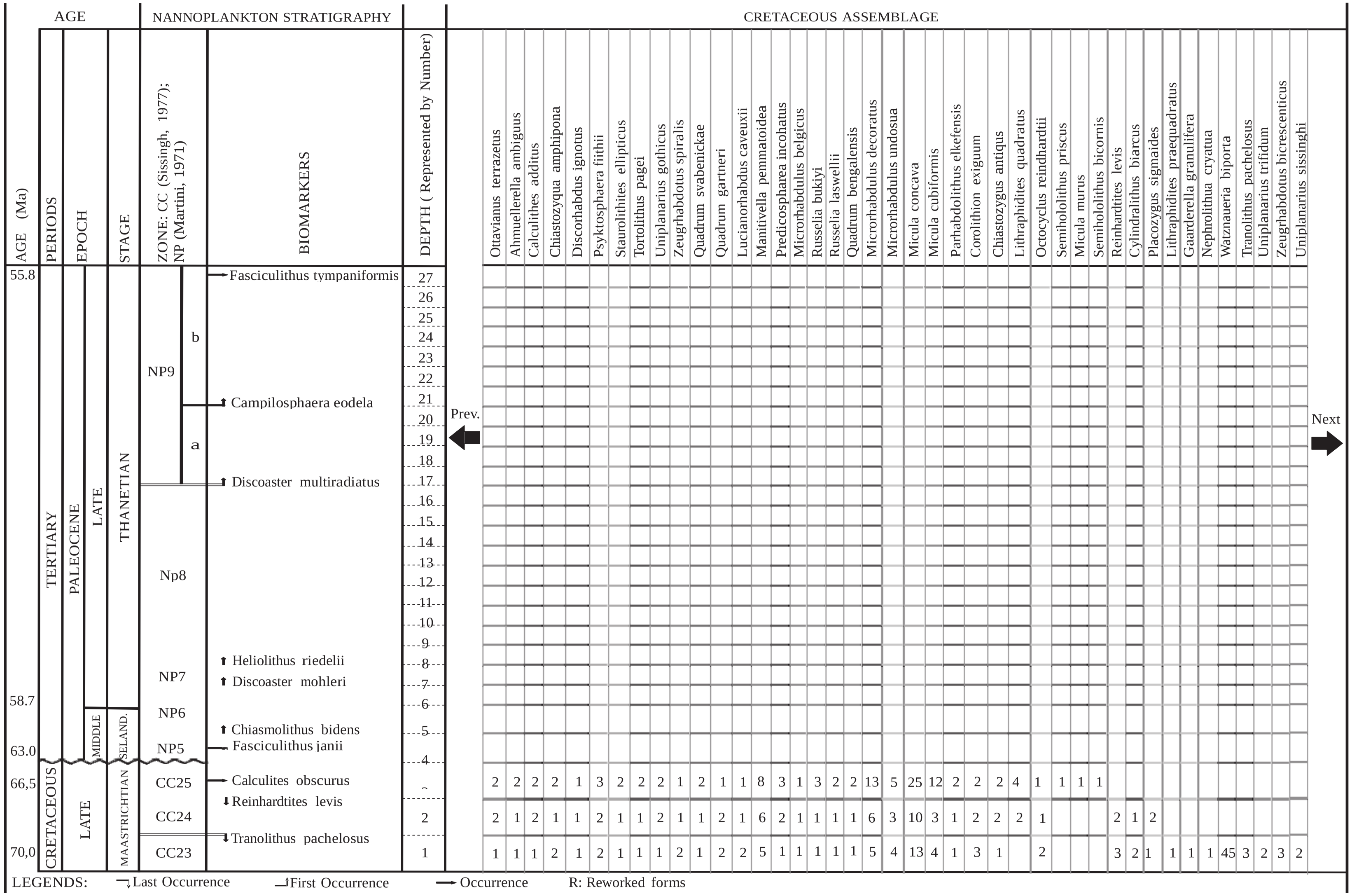
LEGENDS: Last Occurrence; First Occurrence; Occurrence; R: Reworked forms

CRETACEOUS ASSEMBLAGE

	1	2	3	
Prediscosphaera spinosa	7	2	17	
Prediscosphaera stoveri	5	2	13	
Repagulum parvidentatum	2	1	3	
Retecapsa angustiforata	3	1	21	
Retecapsa crenulata	6	1	145	
Retecapsa ficule	3	1	17	
Retecapsa schizobrachiata	2	1	3	
Rhagodiscus angustus	3		4	
Rhagodiscus indistinctus	4	2	7	
Rhagodiscus plebeius	6	2	13	
Rhagodiscus reniformia	5	2	11	
Rhagodiscus splendens	13	5	43	
Staurolithites aurolithites bavus	3	1	5	
Staurolithites flavus	2	1	5	
Staurolithites imbricatus	2	1	2	
Staurolithites integer	1		2	
Staurolithites laaffittei	2	1	3	
Staurolithites zooensia	4	2	5	
Tegumentum stradneri	6	3	12	
Tetrapodorhabdus decorus	3	1	5	
Tranolithus gabulus	5	2	10	
Tranolithus minimus	4	3	12	
Uniplanarius staurophora	10	3	23	
Watznaueria barnesae	1230	103	1430	1R (depth 8), 1R (depth 9), 1R (depth 19), 1R (depth 24)
Watznaueria fossacincta	175	27	23	
Watznaueria ovata	17	13	21	
Watznaueria sp.	4	2	3	
Zeugrhabdotus embergeri	2	1	3	
Zeugrhabdotus erectus	13	5	24	
Zeugrhabdotus praesigmoides	2	1	2	
Zeugrhabdotus trivectis	11	3	25	
Helicolithus anceps	2	2	2	
Broinsonia enormis	1	2	3	
Broinsonia signata	1	1	2	
Ahmuellerella regularis	1	1	1	
Kamptnenus magnificus	1	2	2	
Lithraphidites camiolensis	1	2	2	
Cylindralithua serratus	2	1	1	
Biscutum magnum	1	1	2	

Prev. / Next

Figure 9e. Quantitative Cretaceous nannoplankton distribution chart and biostratigraphy of “P” Well Section.



	Depth 1 (CC23)	Depth 2 (CC24)	Depth ~ (CC25)
Ottavianus terrazetus	1	2	2
Ahmuellerella ambiguus	1	1	2
Calculithes additus	1	2	2
Chiastozyqua amphipona	2	1	2
Discorhabdus ignotus	1	1	1
Psyktosphaera fiithii	2	2	3
Staurolithites ellipticus	1	1	2
Tortolithus pagei	1	1	2
Uniplanarius gothicus	1	2	2
Zeugrhabdotus spiralis	2	1	1
Quadrum svabenickae	1	1	2
Quadrum gartneri	2	2	1
Lucianorhabdus caveuxii	2	1	1
Manitivella pemmatoidea	5	6	8
Predicosphaera incohatus	1	2	3
Microrhabdulus belgicus	1	1	1
Russelia bukiyi	1	1	3
Russelia laswellii	1	1	2
Quadrum bengalensis	1	1	2
Microrhabdulus decoratus	5	6	13
Microrhabdulus undosua	4	3	5
Micula concava	13	10	25
Micula cubiformis	4	3	12
Parhabdolithus elkefensis	1	1	2
Corolithion exiguum	3	2	2
Chiastozygus antiqus	1	2	2
Lithraphidites quadratus		2	4
Octocyclus reindhardtii	2	1	1
Semihololithus priscus			1
Micula murus			1
Semihololithus bicornis			1
Reinhardtites levis	3	2	
Cylindralithus biarcus	2	1	
Placozygus sigmaides	1	2	
Lithraphidites praequadratus	1		
Gaarderella granulifera	1		
Nephrolithua cryatua	1		
Watznaueria biporta	45		
Tranolithus pachelosus	3		
Uniplanarius trifidum	2		
Zeugrhabdotus bicrescenticus	3		
Uniplanarius sissinghi	2		

Figure 9f. Quantitative Cretaceous nannoplankton distribution chart and biostratigraphy of “P” Well Section.

AGE (Ma)	PERIODS	EPOCH	STAGE	ZONE: CC (Sissingh, 1977); NP (Martini, 1971)	BIOMARKERS	DEPTH (Represented by Number)	CRETACEOUS DIVERSITY	CRETACEOUS ABUNDANCE	SURVIVOR DIVERSITY	SURVIVOR ABUNDANCE	PALEOCENE DIVERSITY	PALEOCENE ABUNDANCE	TOTAL DIVERSITY	TOTAL ASSEMBLAGE
55.8	TERTIARY	PALEOCENE LATE	THANETIAN	NP9 b	Fasciculithus tympaniformis	27	0	0	3	3	31	611	34	614
	TERTIARY	PALEOCENE LATE	THANETIAN	NP9 b		26	0	0	1	1	27	168	28	169
	TERTIARY	PALEOCENE LATE	THANETIAN	NP9 b		25	0	0	1	1	30	305	31	306
	TERTIARY	PALEOCENE LATE	THANETIAN	NP9 b		24	0	0	0	0	22	138	22	138
	TERTIARY	PALEOCENE LATE	THANETIAN	NP9 b		23	0	0	4	11	42	1456	46	1467
	TERTIARY	PALEOCENE LATE	THANETIAN	NP9 b		22	0	0	4	10	38	1729	42	1739
	TERTIARY	PALEOCENE LATE	THANETIAN	NP9 b	Campilosphaera eodela	21	0	0	4	16	43	2246	47	2262
	TERTIARY	PALEOCENE LATE	THANETIAN	NP9 a		20	0	0	1	2	13	40	14	42
	TERTIARY	PALEOCENE LATE	THANETIAN	NP9 a		19	0	0	0	0	15	41	15	41
	TERTIARY	PALEOCENE LATE	THANETIAN	NP9 a		18	0	0	3	3	34	326	37	329
	TERTIARY	PALEOCENE LATE	THANETIAN	NP9 a	Discoaster multiradiatus	17	0	0	4	6	35	1294	39	1300
	TERTIARY	PALEOCENE LATE	THANETIAN	Np8		16	0	0	2	2	7	36	9	38
	TERTIARY	PALEOCENE LATE	THANETIAN	Np8		15	0	0	1	1	5	9	6	10
	TERTIARY	PALEOCENE LATE	THANETIAN	Np8		14	0	0	1	1	7	15	8	16
	TERTIARY	PALEOCENE LATE	THANETIAN	Np8		13	0	0	1	1	9	11	10	12
	TERTIARY	PALEOCENE LATE	THANETIAN	Np8		12	0	0	2	2	6	13	8	15
	TERTIARY	PALEOCENE LATE	THANETIAN	Np8		11	0	0	1	1	14	24	15	25
	TERTIARY	PALEOCENE LATE	THANETIAN	Np8		10	0	0	1	1	12	34	13	35
	TERTIARY	PALEOCENE LATE	THANETIAN	Np8		9	0	0	0	0	8	9	8	9
	TERTIARY	PALEOCENE LATE	THANETIAN	Np8	Heliolithus riedelii	8	0	0	4	4	17	40	21	44
	TERTIARY	PALEOCENE LATE	THANETIAN	NP7	Discoaster mohleri	7	0	0	4	5	19	97	23	102
	TERTIARY	PALEOCENE LATE	THANETIAN	NP6		6	0	0	0	0	5	5	5	5
58.7	TERTIARY	PALEOCENE MIDDLE	SELAND.	NP6	Chiasmolithus bidens	5	0	0	1	1	10	14	11	15
63.0	TERTIARY	PALEOCENE MIDDLE	SELAND.	NP5	Fasciculithus janii	4	0	0	3	3	11	20	14	23
66,5	CRETACEOUS	LATE	MAASTRICHTIAN	CC25	Calculites obscurus	3	109	3238	6	21	0	0	115	3259
	CRETACEOUS	LATE	MAASTRICHTIAN	CC24	Reinhardtites levis	2	77	268	3	3	0	0	80	271
70,0	CRETACEOUS	LATE	MAASTRICHTIAN	CC23	Tranolithus pachelosus	1	105	1876	6	8	0	0	111	1884

LEGENDS: Last Occurrence; First Occurrence; Occurrence

Figure 10. Fluctuation of nannoplankton diversity and abundance throughout analyzed section.

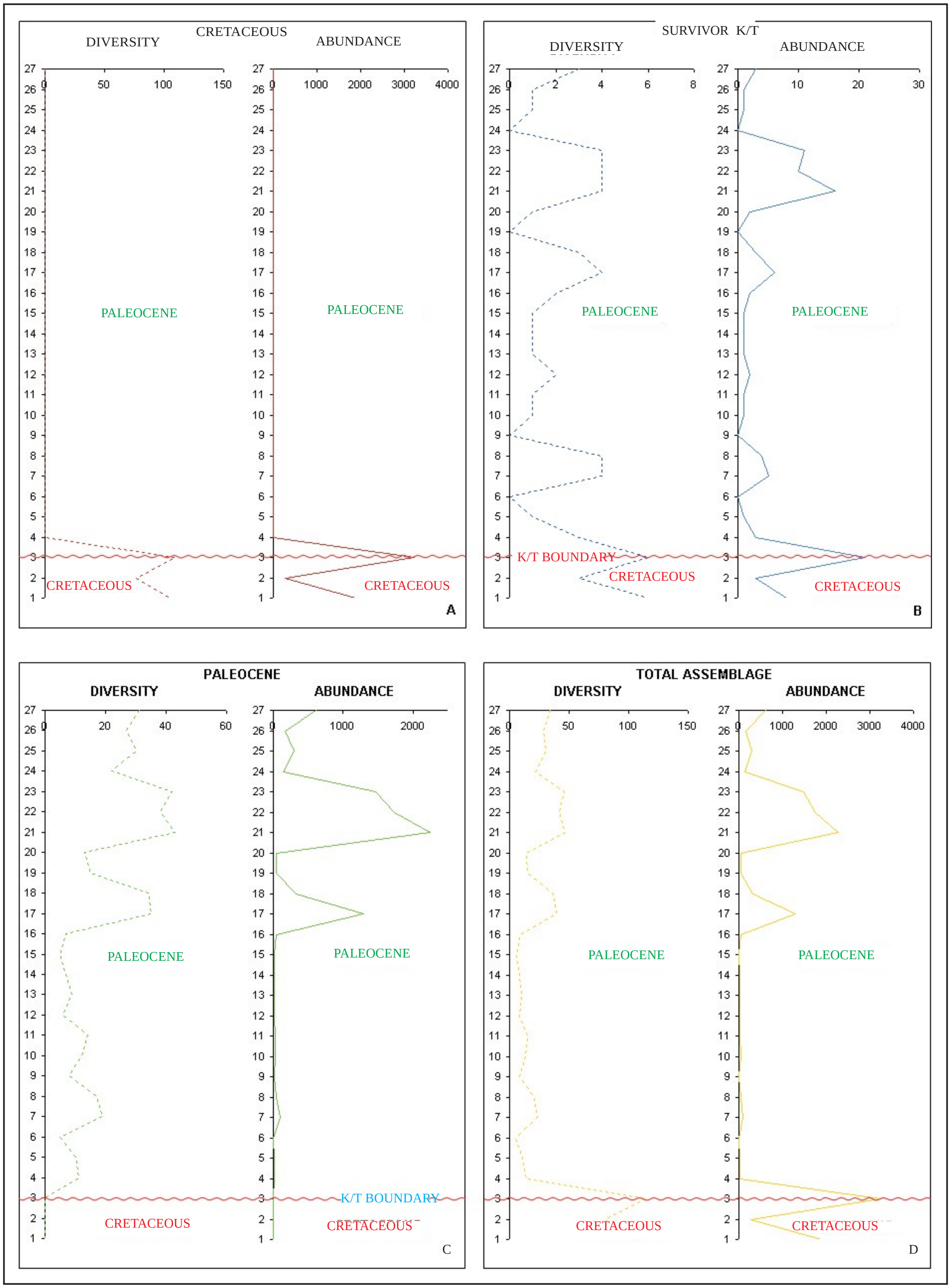


Figure 11. The evolution development graphics, seen in diversity and abundance of Cretaceous, Survivor, and Paleocene nannoplankton assemblages.

(CC25) underwent changing to small and simple plain shapes at the lower part of Paleocene (NP5). In the "P" well section, the Cretaceous nannoplankton assemblages reach 109 species in diversity and 3238 specimens in total assemblage. They are dominated by Genera *Watznaueria, Micula, Arkhangelskiella, Cribrosphaerella, Eiffellithus, Predicosphaera,* and *Retecapsa* (Figure 9D-F). However, none of those dominant species survived in K/T boundary event. A fluctuation in diversity and abundance throughout CC23 to CC25 can be seen clearly.

Survivor Cretaceous species into Tertiary periods consist of 7 species with total assemblage only reach 21 specimen, including *Braarudosphaera bigelowii, Biscutum melaniae, Neocrepidolithus neocrassus, Placozygus sigmoides, Cyclagelosaphaera reinhardtii, Markalius inversus,* and *Scapolithus fossilis* (Figure 9c). They are all minor species within nannoplankton assemblage, but they can survive during K/T boundary event. There is no clearly difference in diversity and abundance fluctuation during Maastrichtian and Paleocene. Their occurrence is always rare to few.

The Paleocene assemblages are characterized by trends of the rising assemblage diversity and abundance of nannoplankton. Generally, the earliest development of new nannoplankton is small and simple coccolith including genera *Praeprinsius Prinsius*, *Neocrepidolithus, Neochiastozygus,* and *Cruciplacolithus,* followed by *Coccolithus, Ericsonia, Toweius, Fasciculithus, Sphenolithus, Ellipsolithus, Chiasmolithus,* and *Zygodiscus* in the Middle, and then closed by *Heliolithus, Discoaster, Helicosphaera, Transversopontis,* and *Lopodolithus.* However, the CC26-NP4 unconformity has removed the small earliest forms. The Paleocene age in this study is initiated by the rare assemblage of genera *Coccolithus, Ericsonia,* medium *Prinsius, Ellipsolithus, Chiasmolithus Toweius, Fasciculithus,* and *Sphenolithus.* The diversity is fluctuated from 5 to 43 species, whilst the abundance from 5 to 2246 specimens is at the upper part of Paleocene. It is dominated by Genera *Toweius, Ericsonia,* and *Coccolithus* (Figure 10).

Rare reworked specimens of Cretaceous nannoplankton are present, including *Micula decussata* and *Watznaueria barnesae* (Figure 9d-e). They can be defined by a different colour (brown) compared with relatively fresh insitu assemblage, etched coccolith, and position in the younger nannoplankton assemblage.

Discussion And Conclusions

The nannoplankton analysis on the "P" well section indicates that 94% of Cretaceous species had been extinguished and remained 6% survivors. A dramatic change in nannoplankton assemblage was found during Paleocene, where the small or simple new species evolved beside the survivors. The peak development of Paleocene nannoplankton evolution was found in the uppermost part of Paleocene.

The nannoplankton succession within "P" well Section is not the real succession within K/T boundary due to the presence of an unconformity ranging from CC26 to NP4. However, it was evident that the nannoplankton assemblages abruptly decreased in diversity and abundance, and mostly changes in species composition.

Fluctuations of diversity and abundance seen in the figure do not completely represent the nannoplankton evolution development (Figure 11). However, that is the final result after the environment of deposition and paleoclimatology are collaborated.

Nannoplankton group dominating at Cretaceous assemblage, are genera *Watznaueria, Micula, Arkhangelskiella, Cribrosphaerella, Eiffellithus, Predicosphaera,* and *Retecapsa* had been extinguished at K/T boundary. It reveals that domination within assemblage is not a guarantee to survive. There is an evidence that they consist of species having restricted tolerance to the environmental change.

Species group consisting of *Braarudosphaera bigelowii, Biscutum melaniae, Neocrepidolithus neocrassus, Placozygus sigmoides, Cyclagelosaphaera reinhardtii, Markalius inversus,* and *Scapolithus fossilis* is the minority in Cretaceous nannoplankton assemblage. However, they survived into the K/T boundary, even *Braarudosphaera bigelowii* and *Scapolithus fossilis* can be found in the present day ocean due to the wide range tolerance to the extreme environmental change during the K/T boundary.

A trend of the rising diversity and abundance can be seen during Late Paleocene. This is the result of the rising temperature at the Late Paleocene after cooling in the Early Paleocene. The lowest sample contains genera *Coccolithus*, *Ericsonia*, medium *Prinsius*, *Ellipsolithus*, *Chiasmolithus Toweius*, *Fasciculithus*, and *Sphenolithus*, whilst the upper part samples contain much more diverse and abundant.

The presence of Cretaceous reworked forms in Paleocene is the evidence of a relative sea level fall in the studied area during Late Maastrichtian to Early Paleocene, that coincided with the Laramide orogeny event in the most area of the world leading to the formation of the Rocky Mountains and the Himalayas.

Acknowledgment–The author would like to thank the Stratigraphy Group, Exploration Department of LEMIGAS Jakarta for the continuing support of nannoplankton researches. This paper is dedicated to researchers who devoutly serve this great country.

References

Alvarez, L.W., Alvarez, W., Asaro, F., and Michel, H.V., 1980. Extraterrestrial cause for the Cretaceous-Tertiary boundary extinction. *Science*, 208, p.1095-1108.

Bagni O. 2007. Brasil Round 4, Santos Basin, Block Definition Superintendency. *http://www.anp.gov.br/brasil-rounds/round4/round4/workshop/restrito/ingles/Santos_ing.pdf*.

Bown, P.R., 1999. *Calcareous Nannoplankton Biostratigraphy*. Kluwer Academic Publishers, p.1-314.

Bramlette, M. N. and Martini, E., 1964. The great change in calcareous nannoplankton fossils between the Maestrichtian and Danian. *Micropaleontology*, 10, p. 291-322.

Geolor's Earth Issues, 2007. Earth's Major Extinctions. *http://www.geolor.com/EarthIssues/Earth_Major_Extinction_geolor.htm* (25 Jan 2007).

HIS, 2007. Information on Brazil, Santos Basin. *http://maps.ihs.com/basin-monitor-ordering-service/latin-america/santos-basin.html* (30 Jan 2007).

Joshua, H., 2007. Play potential in the deepwater Santos basin, Brazil. *Rosenfeld - Yax Balam Inc., Hood, J. F. - Veritas DGC. http://www.pennenergy.com/index/petroleum/display/272394 articles/offshore/volume-66/issue-9/geology-geophysics/play-potential-in-the-deepwater-santos-basin-brazil.html* [30 Jan 2007].

Martini, E., 1971. Standard Tertiary and Quaternary calcareous nannoplankton zonation. *Proceedings, 2nd Planktonic Conference, Roma,* p.739-784.

Okada, H. and Bukry, D., 1980. Supplementary modification and introduction of code numbers to the low latitude coccolith biostratigraphic zonation. *Marine Micropaleontology*, 5 (3), p.321-5.

Perch-Nielsen, K., 1969. Die coccolithen einiger dänischer Maastrichtien - und Danienlokalitäten. *Bulletin of Geological Society, Denmark*, 19, p.51-66.

Perch-Nielsen, K., 1979a. Calcareous nannofossil zonation at the Cretaceous/Tertiary boundary in Denmark. *Proceedings, Cretaceous/Tertiary Boundary Event Symposium, Copenhagen*, 1, p.115-35.

Perch-Nielsen, K., 1979b. Calcareous nannofossils in Cretaceous/Tertiary boundary Sections in Denmark. *Proceedings, Cretaceous/Tertiary Boundary Event Symposium, Copenhagen*, 1, p.120-6.

Perch-Nielsen, K., 1981. Les nannofossiles Calcaires â la Limite Crètacè-Tertiaire près de El Kef, Tunisie. Cah. *Micropaleontol*ogy, 1981, 3, p.25-37.

Perch-Nielsen, K., 1982. Maastrichtian Coccolith in the Danian: Survivors or Reworked "Dead Bodies"? Abstract, *IAS Meeting, Copenhagen*, 122.

Perch-Nielsen, K., 1985. Cenozoic Calcareous Nannofossils, *In*: Bolli, H.M., Saunders, J. B. & Perch-Nielsen, K. (eds.), *Plankton Stratigraphy*. Cambridge University Press, p.329-554.

Percival, S. F. and Fischer, A. G., 1977. Changes in calcareous nannoplankton in the Cretaceous-Tertiary Crisis at Zumaya, Spain. *Evolution Theory*, 2, p.1-35.

Romein, A. J. T., 1977. Calcareous nannofossils from the Cretaceous/Tertiary boundary interval in the Barranco del Gredero (Caravaca, Prov. Murcia, SE Spain). *Proceedings, Koningklijke Nederlandsch Akademie Wetenschappelijk*, 80, p.256-79.

Sharpton, V.L., Dalrymple, G.B., Marin, L.E., Ryder G., Schuraytz, B., and Urrutia, F.J., 1992. New links between Chicxulub impact structure and the Cretaceous-Tertiary boundary. *Nature*, 359, p.819-821

Sissingh, W. 1977. Biostratigraphy of Cretaceous calcareous nannoplankton. *Geologie en Mijnbouw*, 56, p.37-65.

Wikipedia, 2007. Cretaceous-Tertiary extinction event. *http://en.wikipedia.org/wiki/Cretaceous. Tertiary_extinction_event* [25 Jan 2007].